# New data on Plio-Pleistocene dipodine jerboas of the genus *Plioscirtopoda* from Eastern Europe (Mammalia: Dipodidae)

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Remains of dipodine jerboas, *Plioscirtopoda*, are described from the late Pliocene deposits of the Black Sea and Sea of Azov region. Dental morphology of the type species, *P. sepanovi*, is refined. A more primitive species, *P. novorastica* sp. n., is described from the Late Kujalnik deposits of Kryzhanovka 3. The new species differs from *P. stepanovi* in the simpler shape of anterior part of *m*, I layer dimensions and higher numbers of molar roots.

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## Introduction

Plio-Pleistocene dipodine (tridactyl) jerboas were first recorded from Eastern Europe by Gromov & Schevtchenko (1961). The new genus and new species Plioscirtopoda stepanovi was described from the Early Pleistocene Allophaiomys fauna of the upper bed of Kryzhanovka site near Odessa, the Ukraine, Several molars were collected at that time. Gromov & Schevtchenko considered the new form as related to the recent Stylodinus (= Scirtopoda) telum. In late 60's and early 70's V.A. Topachevsky and his group collected and described more material from several localities of the same region. A few molars were added from the type locality and also from the geologically younger Early Pleistocene Allophaiomys faunas of Tarkhankut and Nogaisk. A particularly representative collection of P. stepanovi was described from Tarkhankut, Crimea. The detailed morphological analysis of P. stepanovi by Topachevsky (1973) showed that because of a unique combination of advanced and primitive characters this form cannot be a direct ancestor of recent Stylodipus species, but is rather a specialized independent branch of dipodine ierboas. Since that time, scanty remains of Plioscirtopoda were identified and listed from Early Pleistocene deposits of Don Basin (Kazantseva, 1990; Krasnenkov et al., 1992), Volga area (Sukhov, 1977), Western Siberia (Zazhigin, 1980) and Eastern Siberia (Pre-Baicalian region: Adamenko, 1975, 1977), and from even older, Late Pliocene Khapry faunas (MN17) of Eastern Europe (Topachevsky et al., 1979; Skorik, 1983; Topachevsky & Nesin, 1989), Southern Transurals (Ivakina et al., 1996), and Western Siberia (Zazhigin, 1980). In late 80's and early 90's a large amount of small mammal remains was collected from the stratified Kryzhanovka locality by an expedition of Geological Institute of the USSR Academy of Sciences (Tesakov, 1994; Pevzner et al., 1998). Remains of Plioscirtopoda were obtained from the two successive stratigraphic levels of Kryzhanovka section. Material from Late Pliocene level 3 underlying the type level 4 with P. stepanovi is of special interest. Morphological differences of Plioscirtopoda from level 3 warrant the description of a new species ancestral to P. stepanovi. The present paper deals with dental remains; some data on postreranial morphology of this genus were published by Topachevsky (1973) and Topachevsky & Skorik (1977).

Methods. Dental remains of jerboas were drawn with the use of a camera lucida equiped binocular microscopes MBS-1 and MBS-10. Measurements were taken with the ocular measuring scale of the MBS-10 binocular microscope. The ocular scale was calibrated against standard object scale to remove distor-



Fig. 1. Dental morphology. 1-12, Piloscirtopoda stepanovi I. Gromov & Schevtchenko: 1, 10, 13, m1; 2-4, m2; 2a, m2, anterior side view; 5, m3, 6, 11, M1; 7-8, 12, M2; 9, M3. 1-8, 13 - left; 9, 10-12 - right. 1-9: Kryzhanovka, 4, EMM-25/1-9; 10-12: Tizdat (10 - EMM-29/20; 11-12 - EMM-1065-66). I.S. *Hoisoritopoda novorossica* sq. n., Liventsovka 3.

tions introduced by human vision. All measurements are in millimetres. Scale bars in illustrations correspond to 1 mm. All remains of *Plioscirtopoda* are proserved in Geological Institute of the Russian Academy of Sciences (GIN): coll. EMM-25 (Kryzhanovka 4), EMM-3132 (Kryzhanovka 3), coll. EMM-1029 (Tizdar).

Dental cusps terminology. The basic terminilogy mainly follows that of Shenbrot (1984, 1986) with modifications after Martin (1994). For anterior cusplets in upper molars, anterior and posterior cusps of lower molars, anterocond is used instead of anteroconulid, and hypoconulid instead of posterostylid. Upper case = upper dentition, lower case = lower dentition.

Genus Plioscirtopoda I. Gromov & Schevtchenko, 1961

Diagnosis (emend.). Pm4 absent. Molars mesodont with height in young specimens not exceeding occlusal length. Reentrants deep, not forming enamle listes. Anterior and posterior complications of upper molars present. Cusps alternation well-developed. Anterocomid of m1 half-fused with metaconid. Protoflexid of m2 reduced. Mental foramen shifted almost to middle part of lateral side of mandible.

Plioscirtopoda stepanovi I. Gromov & Schevtchenko, 1961

(Figs 1: 1-12)

Plioscitrapoda stepanovi Gromov & Schevtchenko, 1961: 976-979, fig. 2; Schevtchenko, 1965; 53-54, fig. 354-e; Topachevsky, 1965; 70-74, fig. 16: 1-4, 17; Topachevsky, 1973; 52-74, fig. 10: 1, 11-14, 15: 1; Topachevsky & Skorik, 1977; 69-70; Sukhov, 1977; 88-89, fig. 1: 7-13.

Material examined. Ukrsinc, Odessa Prov., Kryzhanovka 4: 1 m1, 3 m2, 1 m3, 1 M1, 2 M2, 1 M3; Russia, Krasnodar Terr., Tizdar: 1 m1, 2 M2.

Geological age. Early Pleistocene, Early Biharian, small mammal associations dominated by Allophaiomys and lagurines. Diagnosis (emend.). Molars relatively compressed antero-posteriorly. Anterior part of first lower molar flattened, with subequal metaconid and anteroconid.

Description ml. Size:  $2.34 \times 1.85$ . Two round roots. Anterior part slightly biloph because of shallow anterior fold. The fold divides larger rounded metaconid and smaller cusplet, the anteroconid. Anterior tips of metaconid and protoconid at the same level. Metaconid and protoconid the system level.

m2. Size: 1.92 × 1.51,2.12, × 1.85,2.1 × 1.61. Three roots: two anterior and one posterior, flattened root. Posterior, hypoconid part narrower than anterior part. Metaconid and protoconid always widely fused. In one specimen (Fig. 1: 2), shallow second labial reentrant fold (protoficxid) separates protoconid and metaconid portions.

m3. Size:  $1.12 \times 0.98$ . Two rounded roots. Simple c-shaped form with a single deep labial reentrant. A very shallow anterior fold marks a relic of the second labial reentrant (Fig. 1: 5).

M1. Size: about 1.95 x 1.65. Four roots. A fragmentary molar shows meta-, hypo, para-, proto-, and a weekly developed anterocon (Fig. 1: 6). Anterocon points anterolabially. Hypoflexus and metaflexus deep and sightly curving posteriorly (retrovergent). Paraflexus shallow, close to a right angle.

M2. Šize: 1.85 × 1.66, about 2.05 × 1.71. Four roots; lingual roots thicker than labial ones. Occlusal structure similar to that of M1. First labial reentrant (paraflexus) shallow and forms an acute angle. Anterocon points labially. Protocon large. Incipient second lingual flexus present.

M3. Size:  $1.27 \times 1.12$ . Three roots: two anterior and one posterior. Posterior root with signs of fusion of anterior and posterior roots. Ocllusial structure resembles a smaller version of M2 with simplified posterior part.

Hypsodonty. Observed ratio of occlusal length to crown height varies in two anterior molars from 74 to 46 (n = 7).

Notes: The original publication (Gromov & Schevtchenko, 1961) listed Piloscirtopoda and other jerboas within Pilocene Mimomys-Dolomys association of the Kryzhanovka section. However, later Schevtchenko (1965) refined the biostratigraphy of Kryzhanovka samples. Piloscirtopoda remains turned out to be associated with Early Pielsiocene, Allophaiomys small mammal association (Odessa complex). In the present terminology, it corresponds to upper bed of Kryzhanovka (Topachersky et al., 1987; Recovets & Nadachowski, 1995) and Kryzhanovka 4 (Tesakov, 1994), P. stepanovi was based on two molars from Kryzhanovka; the holotype, first upper molar (Gromov & Schevtchenko, 1961; Fig. 2a), and a paratype, second upper molar with absent anteroexternal reentrant fold (ibid., Fig. 2b). However, contrary to recent Stylodipus, M1-2 of Plioscirtopoda never shows complete loss of this fold. Therefore, this specimen seems to be an inverted upper molar, likely M1. When inverted, the molar shows the obtuse anteroexternal fold typical to first upper molars of Plioscirtopoda. Two more molars from upper bed of Kryzhanovka (M1 and M2) were added by Topachevsky (1973). Another close stratigraphic match in Odessa, Zhevakhova Gora 5, yielded three molars: m1, M1, M3 (Topachevsky, 1965, 1973). The new material from the type locality, the upper bed of Kryzhanovka (Fig. 1: 1-9), described in the present paper, makes it possible to emphasize differences between the type form and the well defined Tatkhankut population. The main difference is in the size, the teeth from the Odessa samples (both from Kryzhanovka 4 and Zhevakhova Gora 5) are noticeably larger than those of the Crimean form. However, given the small sample size these differences are not considered statistically meaningful, possibly except for the lower m2, for which the difference is significant at 10% level. For now, it may be noted that the teeth from the two samples of P. stepanovi from the Upper Kujalnik Allophaiomys levels of the Odessa region seem to be larger and more robust as compared to those of other known Early Pleistocene representatives of the species (Nogaisk, Tizdar, Domashkinskie Vershiny). The early Allophaiomys small mammal association of Tizdar (Tesakov, 1998; Pevzner et al., 1998) yielded three molars of P. stepanovi (Fig. 1: 10-12). These molars are close in dimensions to other samples of this species. Dimensions of Tizdar specimens are: 1.95 × 1.42 (m1), 1.71 × 1.32 (M1), 1.78 × 1.50 (M2). The ratio of occlusal length to crown height is rather high: 76 (m1), 117 (M1), 85 (M2). The upper M2 has four roots. The anterior portion of the lower m1 from Tizdar is rounded and lacks a distinct anteroconid

#### Plioscirtopoda novorossica sp. n.

(Figs 1: 13; 2: 1-19)

Plioscirtopoda sp.: Skorik, 1983: 50-51; Topachevsky & Nesin, 1989: 33-34, fig. 9.

Holotype. Isolated lower m1, GIN EMM-31/2 (Fig. 2; 2), Ukraine, Odessa Prov., Kryzhanovka 3.

Paratypes. Kryzhanovka 3: 5 ml, 3 m2, 2 m3, 3 Ml, 4 M2, 1 M3.



Additional material, Liventsovka 3: 1 ml.

Geological age. Late Pliocene, zone MN17a, association of *Mimomys praepliocaenicus* zone dominated by *Borsodia praehungarica*.

Diagnosis. Large Plioscirtopoda with relatively elongated molars. Compact anterior part of first lower molar with metaconid much larger than anteroconid and the anterior fold shifted to the labial side.

Description. Mandible. A fragmentary left mandible from the type locality preserves proximal part with m1-2 and an incomplete incisor (Fig. 3). Masseteric crests on the labial surface of the mandible meet each other below and slightly anterior to the anterior side of m1. They do not reach dorsal portion of the diasterna. Mental foramen placed only slightly dorsal to the midpoint between ventral and dorsal sides of the labial face of lower jaw.

Molars. Occlusal surface flat, formed by distinctly alternating dentine cusps rimmed by narrow enamel bands. Cusps form rounded or triangular prisms with anterior and posterior enamel ridges almost perpendicular to the long axis of the dentition.

m1. Size:  $2.39 \times 1.56$ ,  $2.37 \times 1.66$ ,  $2.2 \times 1.37$ ,  $2.29 \times 1.71$ ,  $2.32 \times 1.61$ ,  $2.34 \times 1.66$ . From two to four roots. Of the four specimens with preserved roots, two have two roots, one has three roots (anterior one is transversely divided into



Fig. 3. Mandible of Plioscirtopoda novorossica sp. n., Kryzhanovka 3, GIN-EMM32/10. a, buccal side; b, lingual side.

two, the posterior root is transversely elongated), and one has four roots (two transversely placed anterior and posterior root pairs). Anterior part of m1 rounded, compact. Metaconid much larger than anteroconid. Anterior tip of metaconid more anterior in position than the anteroconid, the furrow between them shifted to the labial side of the metaconid complex. In one specimen (Fig. 2: 3), anteroconid isolated from the metaconid and represented by a low cingular cusplet. Another young specimen shows the metaconid, protoconid-entoconid, and hypoconid-hypoconulid parts still not connected by wear (Fig. 2: 1). Metaconid-protoconidandentoconid-hypoconid connections not wider than one enamel band width

m2. Size:  $2.29 \times 1.76$ ,  $2.15 \times 1.76$ ,  $2.15 \times 1.76$ ,  $2.15 \times 1.55$ . Two specimens with roots preserved show four roots (one posterior transversely flattende and three anterior ones, two inner roots almost fused). Two lingual and one labial reentrants present. In one specimen, a small second labial reentrant (protoflexid) separates meta- and proto-conids (Fig. 2: 8).

m3. Size: about 1.3 × 1.2. Two roots (anterior and posterior ones). Occlusal surface C-shaped, with one deep labial reentrant.

M1. Size:  $2.2 \times 1.85, 2.39 \times 1.81, 1.95 \times 1.07.$ Four (n = 1) to five (n = 1) roots (postero-labia) root split in two). Two labial and one lingual reentrants; younger specimens have incipient third labial and second lingual reentrants. A small anterocon points anterolabially. Paraflexus (reentrant between proto- and paracone) close to obtuse angle. One young specimen has protoconeparacone and hypocon-metacon pairs still not connected by wear (Fig. 2: 12). Reentrants slightly curving posterioty (refreevregent).

M2. Size:  $1.8 \times 1.32$ ,  $1.98 \times 1.51$ ,  $1.93 \times 1.66$ , ? × 1.76. Of the four specimens, three have four roots and one has five with the posterolingual root split into two. Two labial and one lingual reentrants. Posterior part in younger specimens with shallow incipient reentrants. Small anterocone points labially. Paraflexus shallow and close to an acute angle. Reentrants slightly retrovergent.

M3. Size: 1.51 × 1.22. Two anterior and one posterior roots. Molar biolate. Anterior part formed by broadly fused protocone and paracone. Posterior part formed by fused hypo- and metacones. Anterior and posterior parts connected by very narrow dentine isthmus. Tips of para- and metaflexuses on not meet due to cusp alternation.

Hypsodonty. Ratio of occlusal length to crown height varies in two first molars of lower and upper dentitions from 84 to 55 (n = 14).

Comparison. The molars of P. novorossica sp. n. from Kryzhanovka 3 are larger than those of P stepanovi from the type level of Kryzhanovka 4 and seems to be markedly larger than P stepanovi from Tarkhankut. The difference is significant at 1% level in case of length of lower ml, and at 5% level in case of lower m 2. The new species differs from Pstepanovi in the relatively more elongated first lower molars, more rounded and assymetric metaconid complex, and on the average higher number of molar roots.

Etymology: novorossica from the historical name of the northern coast of the Black Sea.

Notes: A form of Plioscirtopoda geologically older than the type (P. stepanovi) was known for many years on the basis of scanty material from Zhevakhova Gora, bed 11, and middle bed of Kotlovina (Skorik, 1983; Topachevsky & Nesin, 1989). Skorik was first to outline morphological differences of late Pliocene and Early Pleistocene forms. She noted larger dimensions and more elongated molars in the form from a Late Pliocene association of Zhevakhova Gora 11. But Skorik's suggestion that the Late Pliocene from has fewer roots in upper M1 than *P stepanovi* turns fewer roots in upper M1 than *P stepanovi* turns

with bifurcated posterior external roots. The percentage of root number morphotypes cannot be assessed given the present limited material. But P. novorossica sp. n. seems to have a higher number of additional roots in both lower and upper molars. Broad variability of root numbers within samples and gradual decrease of root numbers between fossil and recent members of Eremodipus and Paradipus was described by Shenbrot (1986). According to personal observations, the recent Stylodipus telum also shows a marked variability of root numbers with almost all basic roots capable of splitting into two. A fragmentary first lower molar of Plioscirtopoda from the Late Pliocene fauna of Kotlovina 2 (Topachevsky & Nesin, 1989) matches P. novorossica in the relatively large size and assymetrical anterior complex with small labially shifted anterocon. A single lower m1 of Plioscirtopoda was recovered from level 3 of Liventsovka, a Late Pliocene assemblage geologically younger than Kryzhanovka 3 (coll. of V. Titov, Fig. 1: 13). The specimen shares the structure of anterior complex and molar proportions with P. novorossica and the small size with P. stepanovi. The molar is tentatively assigned to P. novorossica.

### Discussion

Well developed alternation of cusps in lower and upper molars, differentiation of anterior and posterior parts of molars, enamel cutting edges almost perpendicular to the long axis of tooth rows, flattened occlusal surface and structure of proximal part of mandible possibly suggest "microtoid" dental specialization of propalinal jaw movement to increase efficiency of processing green vegetation. Thus, Plioscirtopoda seems to be a specialized steppe genus with predominantly a grass diet. Compared to the Late Pliocene P. novorossica sp. n., the Early Pleistocene P. stepanovi shows reduction in size and number of molar roots, higher differentiation of anterior part of m1, relatively wider occlusal surfaces, and somewhat higher hypsodonty.

*Pitascirappada* is known from Late Pilocene Mimomys faunas (*P. novorassica*) through diversified Allophaiomys faunas of the second part of Early Pleistocene (*P. stepanovi*) in steppe and forest-steppe faunas of Eastern Europe and Western Siberia. The known geological range of those forms extends from about 2.5 to 1.0 Ma. This time range may be even longer because *Piloscirappada* remains were reported from Early-Middle Pilocene Biteke fauna in Western Siberia (Zazhigin, 1980) and from latest Early Pleistocene (Tiraspol) Il'inka fauna in Eastern Europe (Krasnenkov et al., 1992).

The presently known distribution area of Plioscirtopode (Fig. 4) is comparable to that of recent Splodipus (though shifted to the north) and marks Late Pliocene and Early Pliocene steppe and forest steppe zone of western Palearctic. It is interesting that outside the Black Sea area all other finds are situated north of 50th parallel.

A Miocene ancestry of Plioscirtopoda is so far unknown. The Late Turolian Scirtodipus (Savinov, 1970) has plesiomorphic states for most charachters known in Plioscirtopoda. The two genera share the well developed cusps alternation, enamel cutting edges subnormal to long axis of dentition, flat occlusal surfaces, mesodont molars and may form an ancestordescendant pair. The Plioscirtopoda lineage seems to end blindly at the end of Early Pleistocene and there are no direct descendants of this genus in the recent fauna. Among recent jerboas the closest dental match to Plioscirtopoda is Stylodipus. However, molar structure observed in Plioscirtopoda looks more derived in differentiation of elements of occlusal surface as compared to Stylodipus. Each of the three recent species of Stylodipus (Sokolov & Shenbrot, 1987) shares some characters with Plioscirtopoda: S. andrewsi shows reduction of the antero-labial reentrant (protoflexid) of lower m2: S. sungorus has a deep postero-lingual flexid and relics of anteroconid fused with metaconid of m1; all three species, including S. telum, have developed cusp alternation, etc. However, Stylodipus is more hypsodont, has more fused and less pointed cusps, more shallow reentrants, and less flattened occlusal surface than Plioscirtopoda. This suggests that the two genera do not represent a direct phyletic succession and might share a common ancestor (like Scirtodipus) not later than early Pliocene. Most other recent tridactyl jerboas have more or less opposite cusps in dentitions with cuspate occlusal surfaces likely as an adaptation to mixed grain and grass diets. One exception, a specialized desert grass-eating Paradipus, is the most hypsodont modern form with flattened occlusal surfaces. However, this form developed a very different sygmoid dental pattern (Shenbrot, 1986).

According to a cladistic analysis of dental characters in dipodines done by Shenbrot (1992), several genera of tridactyl dipodines form a natural group of closely related forms. According to Shenbrot, they sequentially ramfy from the common ancestral stalk in the fol-



Fig. 4. Geographic distribution of known *Planceriopoda* localities. 1, Kotlovina 2 (Topachevsky & Nessin, 1989); 2. Ocessa region localities: Kryshanovka, Zhevakhova Garo, Tilgal, Luzanovka, etc. (Topachevsky, 1973; Topachevsky et al. 1979; S. S. Loviskovka 3, 6, Middle Dun Itadities Kryshanovka, 1996); P. S. Loviskovka, S. Mogiado En Jones, 1998); S. Loviskovka 3, 6, Middle Dun Itadities Kryshanovka, 1996); P. S. Loviskovka, S. Mogiado En Jones, 1998; S. Loviskovka, 3, 6, Middle Dun Itadities Kryshanovka, 1997; M. S. Loviskov, 1997; J. Tarkatava, 1990; P. J. Parkatava, 1990; P. Mayakava, 1990; P. Mayakavava, 1990; P. Mayakava, 1990; P. Mayakava, 1990; P. Mayakava,

lowing order: Dipus, Stylodipus, Eremodipus, Jaculus. If valid, this scheme likely leaves a place for *Plioscirtopoda* to branch off between Dipus and Stylodipus.

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